

Reliability of the IPCC AR4 Models: A 20th Century Intercomparison of the Freezing/Thawing Index



Motivation

Goals

The annual freezing/thawing index can be used to predict and map permafrost and seasonally ground distribution, frozen active layer and seasonal freeze depths, and has important engineering applications, thereby providing important information on climate variability in cold regions. Most general circulation models (GCMs) do not consider or evaluate permafrost. Therefore, we calculate the freezing/thawing index based on projections of surface air temperatures for the permafrost and seasonally frozen ground regions of the Northern Hemisphere.



"Arctic:" land areas north of 50°N, and "Midlatitudes:" land areas from 20°N-50°N.

We use 5-member ensemble projections of surface air temperatures from the 16 models used in the Intergovernmental Panel on Climate Change (IPPC) Fourth Assessment Report (AR4) to provide an estimate of 21st century freezing/thawing index changes. We make use of four emission scenarios: "commit," "SRESA2," "SRESA1B," and "SRESB1" (see 2001 IPCC report).

An essential consideration in any projections of future climate is the degree to which models can actually reproduce past climate. We evaluate the freezing/thawing index from the IPCC AR4 model simulations' "20th Century Climate in Coupled Models" (20C3M) 5member ensemble runs based on forcing with 20th century historical records, against our Frauenfeld et al. [2007] observed freezing/thawing index data from CRU TS 2.1 (http://nsidc.org/data/ggd649.html).



Figure 2. Freezing (top) and thawing (bottom) index time series for "Arctic," 1901–1999 (left) and 2000-2099 (right). CRU time series (black) is smoothed for ease of comparison with 20C3M. Blue shading corresponds to the 1- σ range among the 16 models. The trend line (red, left) corresponds to the objectively determined 1960s breakpoint in 20C3M.

Results

There is a significant cool bias in 20C3M for the Arctic region (Figure 2, left), manifest as freezing index values that are too high, and thawing index values that are too low. Additionally, the interdecadal variability is underestimated in 20C3M. It should be noted that among the eight regions investigated, this "Arctic" domain nonetheless represents a best-case scenario for agreement between observations and models.

Owing to these potential biases in the model, we next assess the correspondence of the 20C3M simulations against CRU for Northern Hemisphere climatologies. We also calculate trends for the recent decades (since 1960s breakpoint).



Figure 3. The % relative error (RE) between 20C3M and CRU for 1901-1999 (left, freezing index: right, thawing index). Cold season (left): much of the Eurasian Arctic and the Tibetan Plateau is too cold in the models, while midlatitudes are too warmindicated by high REs. Warm season (right): the models are also too cold, though REs tend to be much lower.



Figure 4. Freezing index trends over recent decades (top) are stronger and significant (95%-level) over all land areas in 20C3M (left), while observations (right) indicate significant trends only over certain regions. Similarly, during the warm season, thawing index trends (bottom) are again significant over all land areas, though weaker than observed trends. In general, warm season trends are much weaker than for the cold season, which is correctly captured in 20C3M.

Table 1. Correlations between freezing/thawing index Table 2. Root mean squared error (RMSE) for the from 20C3M and CRU, CAI, and ERA-40; statistically 1958-1998 freezing/thawing index between 20C3M significant Rs (95%-level) are bold; significantly equal and CRU, CAI, and ERA-40. means (1958-1998, t-tests) are italicized.

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Domain	CR	ιU	C.	AI	ERA	\-4 0	Domain	CR	U	CA	١	ERA	-40
Arctic	0.59	0.64	0.61	0.59	0.52	0.54	Arctic	312	46	360	44	395	175
Midlatitudes	0.59	0.74	0.64	0.77	0.50	0.51	Midlatitudes	47	359	46	293	98	588
		Per	mafros	st Regio	ons			Permafrost Regions					
North America	0.33	0.52	0.33	0.51	0.23	0.45	North America	518	66	522	110	254	142
Tibetan Plateau	0.43	0.35	0.63	0.63	0.10	0.18	Tibetan Plateau	724	229	874	582	723	192
Siberia	0.17	0.39	0.27	0.32	0.15	0.27	Siberia	266	155	314	94	761	238
	Seas	onally	Frozer	ı Grou	ıd Regi	ions		Seas	onally	Frozen	Groun	d Regi	ons
North America	0.34	0.58	0.32	0.54	0.32	0.49	North America	67	345	79	199	93	711
Asia	0.51	0.59	0.52	0.64	0.46	0.44	Asia	171	521	112	412	279	776
Russia	0.50	0.52	0.51	0.44	0.48	0.45	Russia	149	286	150	181	184	334

Are these large and significant differences between the models and observations due to shortcomings in CRU? The above tables provide: correlations, t-tests, and RMS errors between 20C3M and the freezing/thawing index from: CRU, Legates and Willmott CAI version 1.02 Tain and ERA-40.

In general, correlations are highest and RMSEs lowest between CRU and 20C3M, i.e. we are not biasing our results by using this particular observational data set; t-tests also indicate that the models have significantly different means in virtually all regions, for all data sets.

Table 3. 1958–1999 trends in freezing (top) and thawing (bottom) index in °Cdays/decade; bold if statistically significant (95%-level). Since 20C3M represents 5-member ensembles of 16 models, the time series are greatly smoothed. Indicated also, therefore, in parentheses, are the normalized trends, to allow for more accurate comparison between regions and data sets.

Domain	CRU		С	AI	ER	A-40	20C3M		
Arctic	-63	(-0.44)	-76	(-0.48)	-52	(-0.39)	-56	(-0.70)	
Midlatitudes	-19	(-0.43)	-28	(-0.52)	-13	(-0.34)	-11	(-0.62)	
			1	Permafrosi	t Regio	ns			
North America	-98	(-0.40)	-113	(-0.43)	-92	(-0.37)	-72	(-0.63)	
Tibetan Plateau	-34	(-0.30)	-236	(-0.63)	3	(0.02)	-52	(-0.65)	
Siberia	-60	(-0.25)	-96	(-0.34)	-55	(-0.21)	-63	(-0.66)	
		S	easona	lly Frozen	Groun	d Region:	5		
North America	-20	(-0.28)	-17	(-0.23)	-14	(-0.25)	-13	(-0.55)	
Asia	-36	(-0.41)	-45	(-0.45)	-26	(-0.35)	-14	(-0.43)	
Russia	-35	(-0.22)	-38	(-0.22)	-27	(-0.18)	-25	(-0.48)	
	CRU		CAI		ERA-40				
Domain	(CRU	С	AI	ER	A-40	20	C3M	
Domain Arctic	25	CRU (0.44)	23	AI (0.42)	ER 18	A-40 (0.32)	20 22	C3M (0.70)	
Domain Arctic Midlatitudes	25 43	CRU (0.44) (0.46)	23 47	(0.42) (0.49)	ER 18 17	A-40 (0.32) (0.20)	20 22 47	C3M (0.70) (0.69)	
Domain Arctic Midlatitudes	25 43	CRU (0.44) (0.46)	23 47	AI (0.42) (0.49) Permafrosi	ER 18 17 t Regio	A-40 (0.32) (0.20)	20 22 47	C3M (0.70) (0.69)	
Domain Arctic Midlatitudes North America	25 43 31	CRU (0.44) (0.46) (0.42)	23 47 1 30	AI (0.42) (0.49) Permafrost (0.42)	ER 18 17 t Regio 19	A-40 (0.32) (0.20) ons (0.29)	20 22 47 14	C3M (0.70) (0.69) (0.59)	
Domain Arctic Midlatitudes North America Tibetan Plateau	25 43 31 5	(0.44) (0.46) (0.42) (0.10)	23 47 1 30 192	(0.42) (0.49) Permafrost (0.42) (0.58)	ER 18 17 t Regio 19 0	A-40 (0.32) (0.20) ms (0.29) (-0.01)	20 22 47 14 17	C3M (0.70) (0.69) (0.59) (0.60)	
Domain Arctic Midlatitudes North America Tibetan Plateau Siberia	25 43 31 5 15	(0.44) (0.46) (0.42) (0.10) (0.24)	23 47 30 192 17	AI (0.42) (0.49) Permafrost (0.42) (0.58) (0.27)	ER 18 17 t Regio 19 0 10	A-40 (0.32) (0.20) ms (0.29) (-0.01) (0.15)	20 22 47 14 17 16	C3M (0.70) (0.69) (0.59) (0.60) (0.58)	
Domain Arctic Midlatitudes North America Tibetan Plateau Siberia	25 43 31 5 15	(0.44) (0.44) (0.46) (0.42) (0.10) (0.24) <i>S</i>	23 47 30 192 17 Seasona	AI (0.42) (0.49) Permafrost (0.42) (0.58) (0.27) Illy Frozen	ER 18 17 t Regio 19 0 10 Groun	A-40 (0.32) (0.20) ms (0.29) (-0.01) (0.15) md Regions	20 22 47 14 17 16	C3M (0.70) (0.69) (0.59) (0.60) (0.58)	
Domain Arctic Midlatitudes North America Tibetan Plateau Siberia North America	25 43 31 5 15 34	(0.44) (0.46) (0.42) (0.10) (0.24) (0.37)	C 23 47 1 30 192 17 5easona 31	AI (0.42) (0.49) Permafrost (0.42) (0.58) (0.27) Illy Frozen (0.33)	ER 18 17 t Regio 19 0 10 Groun 19	A-40 (0.32) (0.20) ms (0.29) (-0.01) (0.15) ad Regions (0.22)	20 22 47 14 17 16 5 47	C3M (0.70) (0.69) (0.59) (0.60) (0.58) (0.67)	
Domain Arctic Midlatitudes North America Tibetan Plateau Siberia North America Asia	25 43 31 5 15 34 34	(0.44) (0.46) (0.42) (0.10) (0.24) S (0.37) (0.34)	C 23 47 1 30 192 17 <i>Seasona</i> 31 42	AI (0.42) (0.49) Permafrost (0.42) (0.58) (0.27) Illy Frozen (0.33) (0.42)	ER 18 17 t Regio 19 0 10 Groun 19 22	A-40 (0.32) (0.20) ms (0.29) (-0.01) (0.15) md Regions (0.22) (0.21)	20 22 47 14 17 16 5 47 32	C3M (0.70) (0.69) (0.59) (0.60) (0.58) (0.67) (0.63)	

Conclusions

We note some significant differences between models and observations over the 20th century. The models are too cold (higher freezing indices) in high latitudes/altitudes during the cold season, and too warm in middle latitudes. During the warm season, the models are too cold (lower thawing indices) over most of the Northern Hemisphere. Trends over recent decades are greatly overestimated by the models as well, resulting in too much warming that is too widespread. While we do show the freezing/thawing index for the Arctic for various emission scenarios over the 21st century in Figure 2, we urge caution given the potential model shortcomings over the 20th century.

Trends suggest IPCC models are statistically significantly warming for all Northern Hemisphere permafrost and seasonally frozen ground regions. However, the three observational data sets indicate distinct regional differences. suggesting that the models may be overestimating historical warming in the 20th century. Furthermore. when calculating the normalized trends (accounting for lower standard deviation in the smoother model time series), the 20C3M model trends are much greater in all instances, almost twice the the magnitude of observed trends in many cases.

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