# Probabilistic evaluations of a cloud system model using ARM observations Pete Henderson & Robert Pincus. CIRES/NOAA, University of Colorado, Boulder PeterH@Colorado.Edu, Robert.Pincus@Colorado.Edu

#### Introduction

- 3-year runs of a CSRM<sup>1</sup> in quasi-MMF<sup>2</sup> configurations are made over the SGP<sup>3</sup> site (1999-2001) using ARM's observationally-constrained forcing data [1].
- To constrain our evaluation of cloud, winds & temperature are nudged back to observed values every 2 & 24 hours.
- Evaluations of model cloud-occurrence are then made against ARM'S ARSCL<sup>4</sup> data [2].



**FIGURE 1.** (left) Snapshot of the hourly mixing ratios of a single column of the model's domain. (right) Corresponding total reflectivity simulated by Quick-Beam [2].

## Method

- To make observed and modeled cloud more comparable, ARSCL's sensitivity thresholds are used to define cloud occurrence in the model (Figs.1–2).
- Using the model's mixing ratios of cloud liquid and ice, rain, snow and graupel; we simulate observations made by ARSCL's radar using Quick-Beam [3], and ARSCL's lidar with an approximate calculation of the extinction coefficients,  $k_{ext, liq}$  and  $k_{ext, ice}$ . Effective radii,  $r_{\rho}$  from the CAM<sup>5</sup> are used to find the droplet number concentration, N from which k<sub>ext</sub> follows, assuming a scattering efficiency of 2.

**FIGURE 2.** Probability distributions for two model runs (left & right), of simulated radar reflectivities (top-row) relative to the radar's sensitivity threshold (blue) and approximated lidar extinction coefficient (bottom-row),



- The "lidar-proxy" defines cloud when  $k_{ext} > 0$ , unless optical depth > 2, in which case the beam is assumed to be fully extinguished and the clear-cloud decision is made by the simulated radar component alone, which defines an event as cloudy if it exceedes the the sensitivity threshold, shown in Figs.1–2.
- At each time-step, the model's probability of cloud (PoC) is determined from the binary decision.
- Although simple, the lidar-proxy does improve estmations of ARSCL's cloud-based during precipitation.

### **Provisional results**

- ARSCL for z < (>) 7 km.



#### **References:**

[1] Xie, S., R. T. Cederwall and M. Zhang, 2004: Developing long-term SCM/CSRM forcing data using NWP products constrained by surface and TOA observations. J. of Geophys. Res., 109, D01104.

[2] Clothiaux, E. E., T. P. Ackerman, G. G. Mace, K. P. Moran, R. T. Marchand, M. A. Miller, and B. E. Martner, Objective Determination of Cloud Heights and Radar Reflectivities Using a Combination of Active Remote Sensors at the ARM CART Sites, 2000: J. Appl. Meteorology, 39, 645–665.

[3] Haynes, J. M., R. T. Marchand, Z. Luo, A. Bodas-Salcedo, and G. L. Stephens, 2007: A multi-purpose radar simulation package: Quick-Beam. Bull. Amer. Meteor. Soc., 88, 1723–1727.

FOOTNOTES: 1. Cloud System Resolving Model. 2. Multiscale Modeling Framework. 3. Southern Great Plains. 4. Active Remotely-Sensed Cloud Locations. 5. Community Atmospheric Model

• Fig.3 shows that both runs reproduce the overall shape of ARSCL's mean profile.

• Differences in the local maxima of mean PoC, near 1 km and 8–10 km, correspond to those in the pdfs of Fig.2.

• IPHOC is in better(worse) agreement with

mean cloud ocurrence in ARSCL.