

Weather Research and Forecasting Model

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Outline

- What does WRF model do?
- WRF Standard Initialization
- WRF Dynamics
 - Conservation Equations
 - Grid staggering
 - Time integration
 - Boundary conditions

WRF Physics Options

- Turbulence/Diffusion
- Radiation
- Surface
- PBL
- Cumulus parameterization
- Microphysics

Testing, Verification, and Computational Efficiency

What does WRF model do?

• Developed by:

NCAR/MMM and NOAA/FSL with partnership at NCEP, AFWA, FAA, NRL and collaborations with other universities.

- Latest version: WRF version 2.0
- Developed for:
 - Idealized simulations (convection, baroclinic waves, large eddy simulations)
 - Parameterization research
 - Data assimilation research
 - Forecast research
 - Real-time NWP
 - Coupled-model applications

WRF Modeling System Flow Chart (for WRFV2)



WRF Standard Initialization

- Provides all required initial and time-varying boundary conditions
- De-grib GRIB files for meteorological data
- Provides method to define and localize WRF domain, nests, and subnests
- Produces terrain, landuse, etc on domain:
 - USGS 30 sec (~1km) topography
 - USGS 24 category landuse
 - WMO/FAO 16 category 2-layer soil types
 - Annual mean deep soil temperatures
 - Monthly greenness fraction
 - Albedo
 - Terrain slope index
 - Max Snow Albedo



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WRF Standard	Sources Script	t				
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WRF Dynamics

- Terrain Representation
- Vertical Coordinate
- Grid Staggering
- Time integration scheme
- Conservation Equations
- Advection Scheme
- Boundary conditions

Terrain Representation

 Lower boundary condition for geopotential specifies the terrain elevation

$$\frac{\partial \phi}{\partial t} + u \frac{\partial \phi}{\partial x} + v \frac{\partial \phi}{\partial y} + \omega \frac{\partial \phi}{\partial \eta} = gw$$



Vertical Coordinate

Hydrostatic pressure

$$\pi \qquad \eta = \frac{(\pi - \pi_t)}{\mu}, \qquad \mu = \pi_s - \pi_t$$

Vertical resolution set by user in WRFSI (default: 31 levels)
Can choose option for vertical interpolation (log, linear, square root)

Grid Staggering

C grid spacing

Time-split Leap Frog and Runge-Kutta scheme

- 3rd order RK generally stable using timestep twice as large in leapfrog model
- Courant number limited to $C_r = U\Delta t / \Delta x < 1.73$







Wicker and Skamarock, 2002 MWR

- RK3 method excellent scheme for integrating the compressible equations and is ideal candidate for NWP
- RK3 best combination of accuracy and simplicity

TABLE 1. Maximum stable Courant number for one-dimensional linear advection. Here, U indicates the scheme is unstable.

	Spatial order			
Time scheme	3rd	4th	5th	6th
Leapfrog RK2 RK3	U 0.88 1.61	0.72 U 1.26	U 0.30 1.42	0.62 U 1.08





Phase and Amplitude errors for LF and RK3

Advection equation analysis

 $\phi_t = U\phi_x$

5th and 6th order upwind-biased and centered schemes. Analysis for 4∆x wave.



Phase and Amplitude errors for LF and RK3

Oscillation equation analysis

 $\phi_t = ik\phi$



Conservation Equations Mass Coordinate (Flux)

$$\frac{\partial U}{\partial t} + \mu \alpha \frac{\partial p}{\partial x} + \frac{\partial p}{\partial \eta} \frac{\partial \phi}{\partial x} = -\frac{\partial U u}{\partial x} - \frac{\partial \Omega u}{\partial \eta}$$
$$\frac{\partial W}{\partial t} + g \left(\mu - \frac{\partial p}{\partial \eta} \right) = -\frac{\partial U w}{\partial x} - \frac{\partial \Omega w}{\partial \eta}$$
$$\frac{\partial \Theta}{\partial t} + \frac{\partial U \theta}{\partial x} + \frac{\partial \Omega \theta}{\partial \eta} = \mu Q$$
$$\frac{\partial \mu}{\partial t} + \frac{\partial U}{\partial x} + \frac{\partial \Omega}{\partial \eta} = 0$$
geopotential ($\phi = gz$) $\frac{d\phi}{dt} = gw$
Hydrostatic pressure Ideal Gas Law
 $\frac{\phi}{p} = -\mu \alpha, \qquad p = \left(\frac{R\theta}{p,\alpha}\right)^r, \quad \Omega = \mu \eta$

 $p_0 \alpha$

 $\partial \phi$

дŋ

Momentum

Heat Continuity

Diagnostic Relations

Conservation Equation Height Coordinates (Flux)

$$\begin{split} U &= \rho u, \quad V = \rho v, \quad W = \rho w, \quad \Theta = \rho \theta \\ \frac{\partial U}{\partial t} + \gamma R \pi \frac{\partial \Theta}{\partial x} - f V &= -\frac{\partial U u}{\partial x} - \frac{\partial W u}{\partial z} \\ \frac{\partial W}{\partial t} + \gamma R \pi \frac{\partial \Theta}{\partial z} + g \rho &= -\frac{\partial U w}{\partial x} - \frac{\partial W w}{\partial z} \\ \frac{\partial \Theta}{\partial t} + \frac{\partial U \theta}{\partial x} + \frac{\partial W \theta}{\partial z} = \rho Q \\ \frac{\partial \rho}{\partial t} + \frac{\partial U}{\partial x} + \frac{\partial W}{\partial z} = 0 \\ \gamma R \pi \nabla \Theta &= c_p \Theta \nabla \pi = \nabla p \end{split}$$

Conservative Variables

Momentum

Heat

Continuity

Pressure terms related to Θ

Moist Equations

$$\begin{aligned} \frac{\partial U}{\partial t} + \alpha \mu_{d} \frac{\partial p}{\partial x} + \frac{\alpha}{\alpha_{d}} \frac{\partial p}{\partial \eta} \frac{\partial \phi}{\partial x} &= -\frac{\partial U u}{\partial x} - \frac{\partial \Omega u}{\partial \eta} \\ \frac{\partial W}{\partial t} + g \left(\mu_{d} - \frac{\alpha}{\alpha_{d}} \frac{\partial p}{\partial \eta} \right) = -\frac{\partial U w}{\partial x} - \frac{\partial \Omega w}{\partial \eta} \\ \frac{\partial \mu_{d}}{\partial t} + \frac{\partial U}{\partial x} + \frac{\partial \Omega}{\partial \eta} = 0 \\ \frac{\partial (\mu_{d} q_{v,l})}{\partial t} + \frac{\partial U q_{v,l}}{\partial x} + \frac{\partial \Omega q_{v,l}}{\partial \eta} = \mu_{d} Q_{v,l} \end{aligned}$$

$$\begin{aligned} \mathbf{Hydrostatic Pressure} \quad \mathbf{Ideal Gas Law} \\ \frac{\partial \phi}{\partial \eta} = -\alpha_{d} \mu_{d}, \quad p = \left(\frac{R\Theta}{p_{o} \mu_{d} \alpha_{v}}\right)^{r} \end{aligned}$$

Momentum

Continuity

Heat

Diagnostic Relations

Advection

• 2nd, 3rd, 4th, Example: 5th order scheme 5th,and 6th order centered and upwind biased schemes

$$\frac{\partial (U\phi)}{\partial x} = \frac{1}{\Delta x} \left(F_{i+\frac{1}{2}}(U\phi) - F_{i-\frac{1}{2}}(U\phi) \right)$$

where

$$\begin{split} F_{i-\frac{1}{2}}(U\phi) &= U_{i-\frac{1}{2}} \left\{ \frac{37}{60} (\phi_i + \phi_{i-1}) - \frac{2}{15} (\phi_{i+1} + \phi_{i-2}) + \frac{1}{60} (\phi_{i+2} + \phi_{i-3}) \right\} \\ &- sign(1,U) \frac{1}{60} \left\{ (\phi_{i+2} - \phi_{i-3}) - 5 (\phi_{i+1} - \phi_{i-2}) + 10 (\phi_i - \phi_{i-1}) \right\} \end{split}$$

$$\Delta t \frac{\delta (U\phi)}{\Delta x} \Big|_{5th} = \Delta t \frac{\delta (U\phi)}{\Delta x} \Big|_{6th} \qquad \text{For constant U}$$
$$- \frac{\left| \frac{U\Delta t}{\Delta x} \right|_{6th}}{\left| \frac{1}{60} \left(-\phi_{i-3} + 6\phi_{i-2} - 15\phi_{i-1} + 20\phi_i - 15\phi_{i+1} + 6\phi_{i+2} - \phi_{i+3} \right)}{\frac{Cr}{60} \frac{\partial^6 \phi}{\partial x^6} + H.OT}$$

Boundary Conditions

Тор

- 1. Constant pressure
- 2. Absorbing upper layer (increased horizontal diffusion) Bottom
 - 1. Free Slip
- 2. Variety BL implementations on surface drag and fluxes Lateral

1. Specified

- Open (perturbations can pass into/out of model domain)
 Symmetric
 - Periodic (values of dependent variables are assumed identically equal to values of another boundary)
- 5. Nested

WRF Physics

- Subgrid Eddy diffusionPBL
- Cumulus parameterization
- Radiation
- Microphysics
- Surface

Parameterizations Interactions





Parameterizations Interactions

 It appears that the physic options are done within each RK loop EXCEPT the microphysics

Microphysics:
Heat/moisture tendencies
Microphysics rates
Surface rainfall

Begin time step



End time step

Physics	Options	Comments
Subgrid eddy diffusion	 Constant diffusion Level 2.5 TKE Stress/deformation Smagorinsky 2D Horizontal Smagorinsky 	 Const khdif and kvdif Based on K Based on horiz wind for horiz diffusion only
Boundary Layer* * Includes surface similarity theory	 YSU MRF Mellor-Yamada-Janjic (Eta) 	 Non-local K mixing in dry convective BL, vertical diffusion depends on Ri 1.5 order, level 2.5, TKE prediction with Local K vertical mixing
Convective parameter.	 new Kain-Fritsch Betts-Miller-Janjic (Eta) Grell Ensemble 	 shallow convect, mass flux up/down draft no explicit up/downdraft multiple closure and parameter, explicit up/downdraft

Physics	Options	Comments	
Longwave Radiation	• RRTM (MM5)	Spectral scheme with K distribution and a look up table	
	• Eta (GFDL)	 Spectral scheme from global model used in Eta 	
Shortwave	• Dudhia (MM5)	 Simple downward calculation, clear scattering 	
Radiation	Goddard	•Spectral method	
	• Eta (GFDL)	 Used in Eta, ozone effects and interacts with clouds 	
Land-	 5 layer thermal diffusion 	 layers 1,2,4,8,and 16 cm thick 	
Surface	NOAH Land Surface	 soil temp and moisture 4 layers 	
	RUC Land Surface	 soil temp and moisture 6 layers 	
Micro-	• Kessler	 warm rain, no ice, idealized 	
physics	• Lin et al.	 5 class including graupel 	
	• WSM3	• 3 class with ice, ice processes	
	• WSM5	• 5 class with ice, supercooled H_20	
	• Eta (Ferrier)	one prognostic total condensate	
	• WSM6	6 class with graupel	

Testing and Verification

 Simulations run to target specific facets of research or forecasting.



Testing and Verification



Testing and Verification

a = 1 km, dx = 200 m

a = 100 km, dx = 20 km



Column Maximum Reflectivity 00 UTC 16 JUNE 2002

a) National Radar Composite



c) 10KM Resolution, LIN, KAIN-FRITSCH



b) 4 KM Resolution, LIN Microphysics



d) 10KM Resolution, NCEP3, KAIN-FRITSCH



Verification

Weisman et al. – Although the 4 km is typically minimum grid resolution, there was significant improvement in

> representing the system scale structure for larger convective systems

Isolated convective outbreaks were not as well represented
 During International H₂0 project (IHOP)

Verification

Baldwind and Wandishin

- Found that WRF reproduced the observed spectra much better than higher resolution Eta
 10 km WRF model forecast maintains the variance in precip field down to at least 4 times grid spacing (40km)
 - grid spacing (40km) Variance of Eta drops off sooner and at greater than 10 times grid spacing



3 hr accumulated precipitation valid at 18Z from 12Z 4 June 2002 model run.



Computational Efficiency

Michalakes et al.

- WRF is more costly in terms of time-per-time step
- BUT the RK3 allows for considerably longer time step (200sec WRF vs. 81sec MM5)
- Time-to-solution performance for WRF slightly better than MM5
 - Authors feel this will improve with tuning and optimization



36km resolution on a 136 x112 x 33 grid

Summary of WRF

- Fully compressible, non-hydrostatic (with hydrostatic option)
- Eulerian mass/height based terrain following coordinate
- Arakawa C staggering
- Runge-Kutta time integration scheme
- Higher order advections
- Scalar-conserving
- Complete Coriolis and curvature
 Two-way and one-way nesting
 Lateral boundary conditions for ideal or real data
 Full physics options

Questions?

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References

Michalakes et al.: Development of a Next-Generation Regional Weather Research and Forecast Model

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Shamarock et al., 2001: Prototypes for the WRF Model http://www.mmm.ucar.edu/individual/skamarock/meso2001pp_wcs.pdf

Weisman et al. 2002, : Preliminary Results from 4km Explicit Convective Forecasts Using the WRF Model. (Preprint) AMS 19th Conf. on Weather Analysis and Forecasting and 15th Conf. on Numerical Weather Prediction. Aug. 12^{th.}

Wicker L.J. and W. Shamarock, 2002: Time-Splitting Methods for Elastic Models Using Forward Time Schemes. *Mon. Wea. Rev.,* Vol. 130, pg. 2088-2097.

VRF model Users Web site: http://www.wrf-model.org