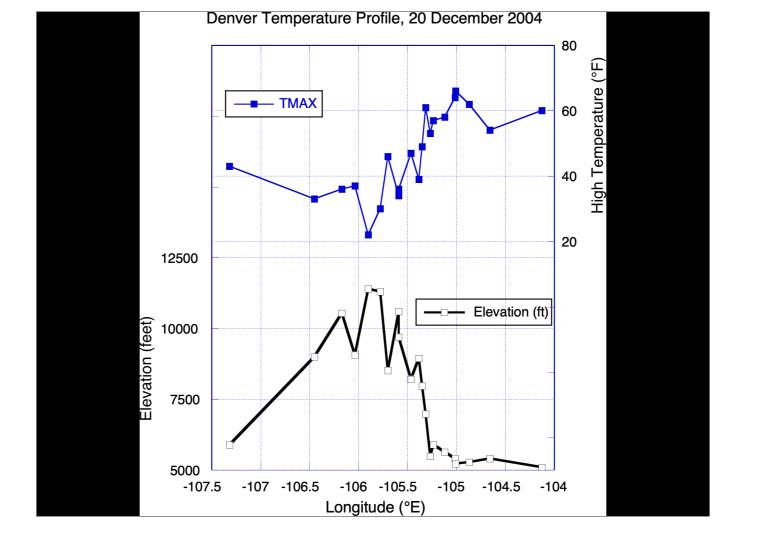
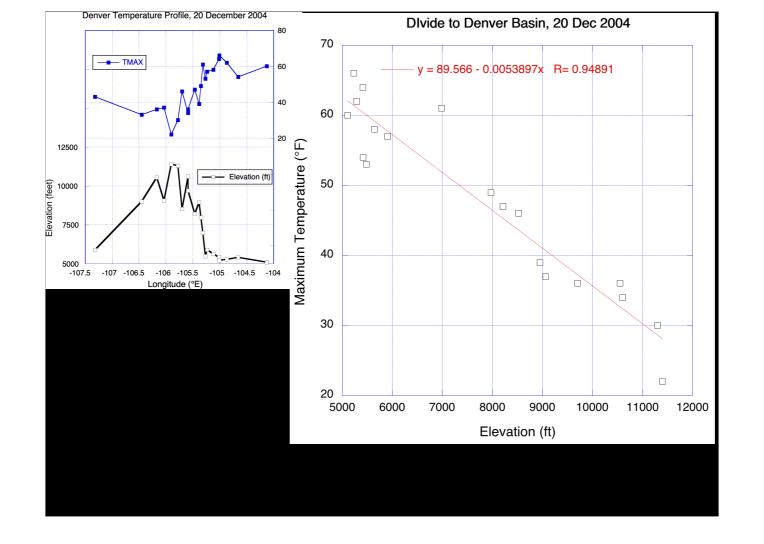
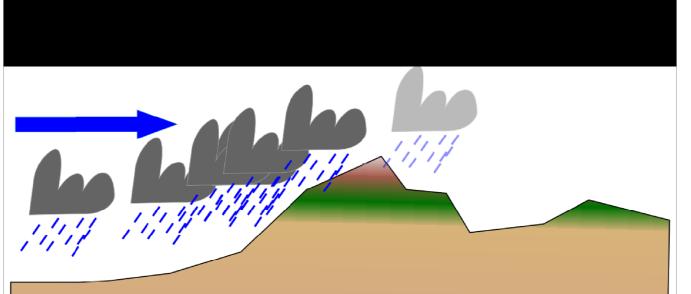
• What varies with elevation?





What varies with elevation?

• Temperature (dry lapse rate 9.8°C/km)

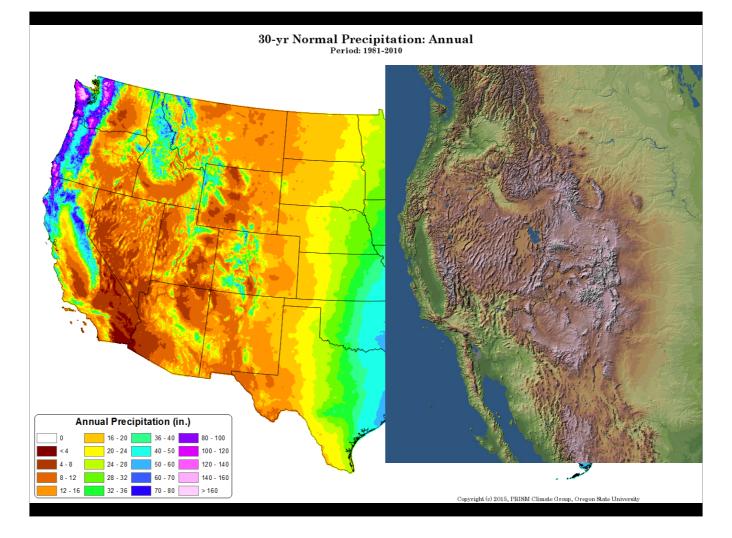


How does temperature vary in the rain?

Not trivial to answer, but making rain is exothermic, so moist lapse rate more like 5°C/km

What varies with elevation?

- Temperature (dry lapse rate 9.8°C/km, moist ~5°/km)
- Precipitation



Locally well correlated, but at broader scales more going on

What varies with elevation?

- Temperature (dry lapse rate 9.8°C/km, moist ~5°/km)
- Precipitation
- Radiation
- Air pressure
- Slopes (derivative)

Gravity, too, but nobody has found that to be a tool...

What varies with elevation?

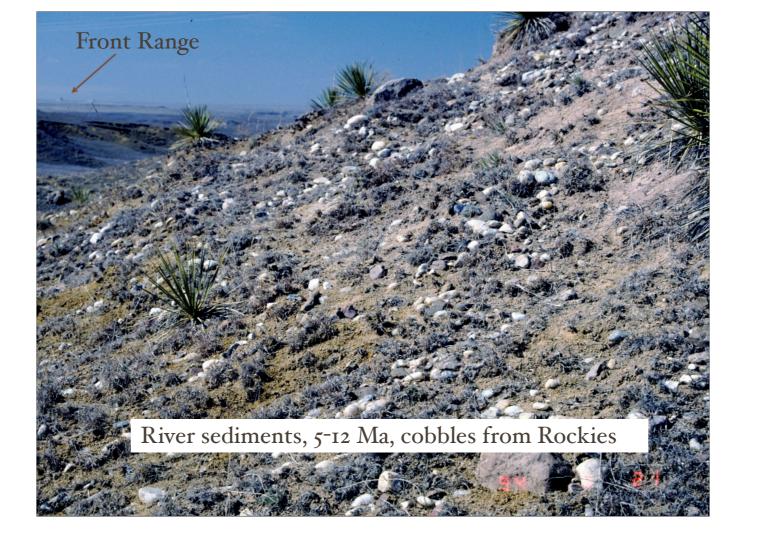
- Temperature (dry lapse rate 9.8°C/km, moist ~5°/km)
 - Paleobotany
 - Clumped isotopes
- Precipitation
 - Paleobotany
 - Oxygen and hydrogen isotopes (Rayleigh distillation)
- Radiation
 - Cosmogenic isotopes
- Air pressure
 - Plant stomatal density
 - Bubbles in lava
- Slopes (derivative)
 - Erosion, deposition

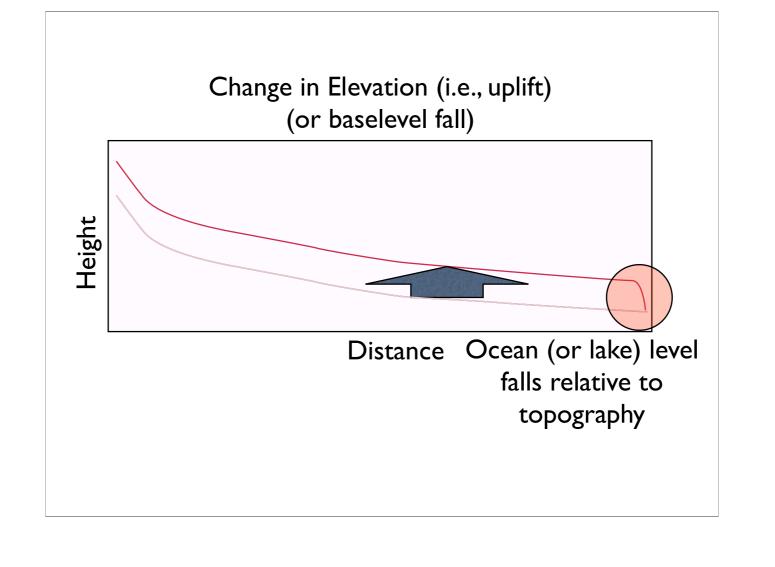
You'd think with all these that this would be open and shut...but all of them have issues.

What varies with elevation?

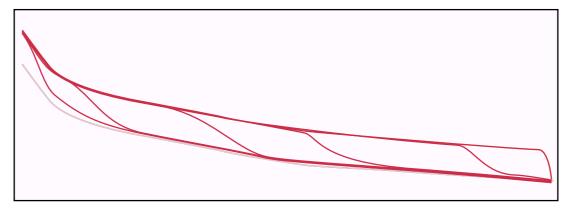
- Temperature (dry lapse rate 9.8°C/km, moist ~5°/km)
 - Paleobotany
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River erodes rapidly where gradient is steep...



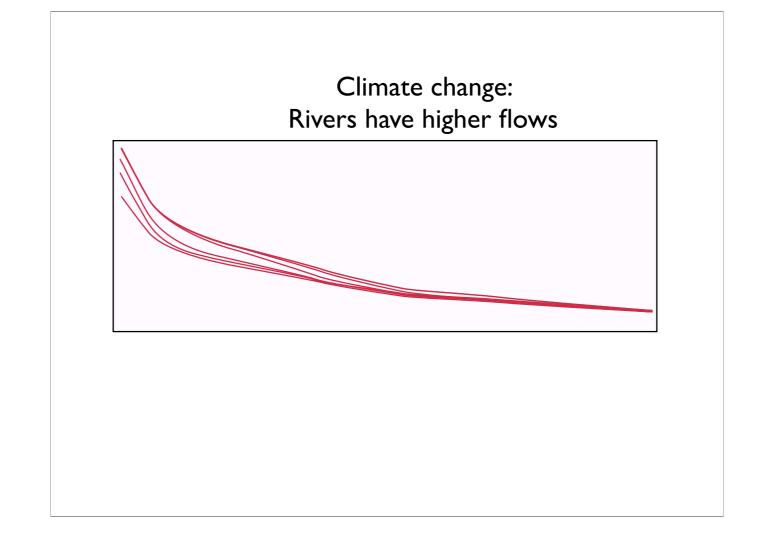
leading to restoration of original river profile but, for quite some time, leaving remnants of old topography (peaks, canyons)

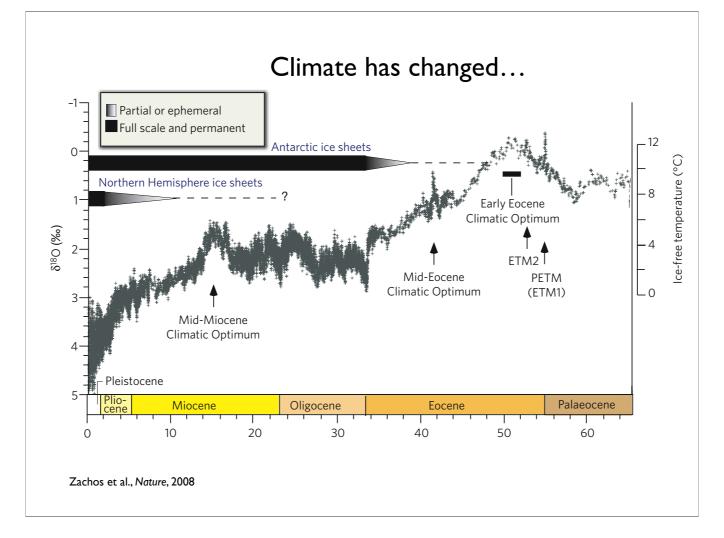
What makes streams erode?

- I. Faster water flows, more it can lift and carry.
- 2. More water flowing, more it can lift and carry.

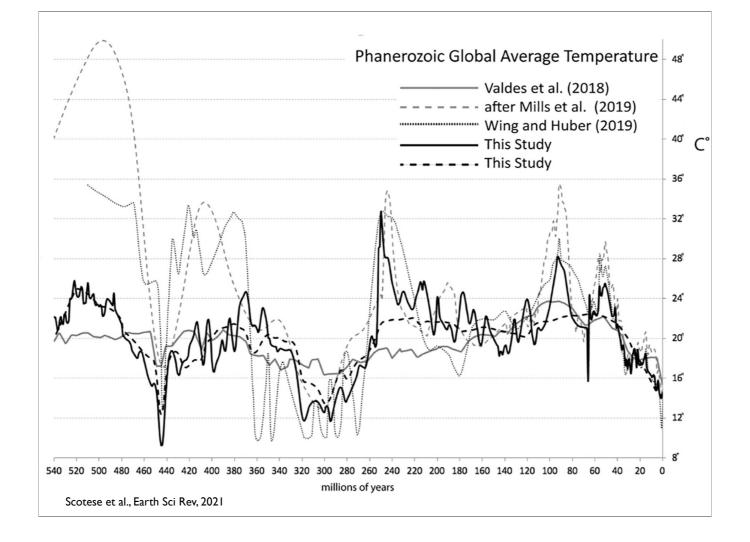
$$E = k \left(\frac{\rho g QS}{W} - w_0 \right)$$

Erosion increases as the flow rate *Q* (volume/unit time) increases, as *W* (channel width) decreases, and as *S* (slope) increases. Other terms constants.

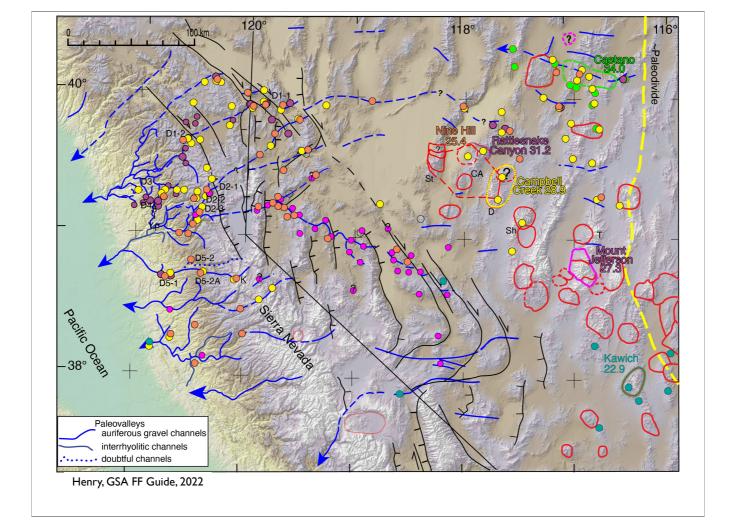


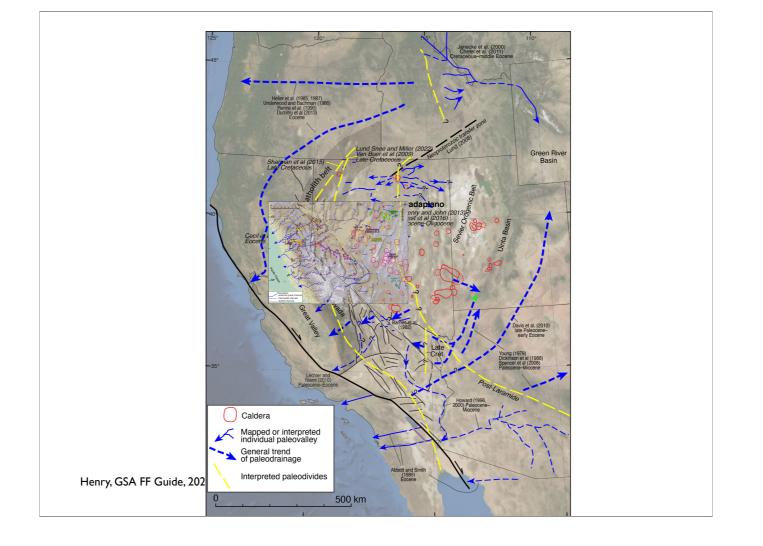


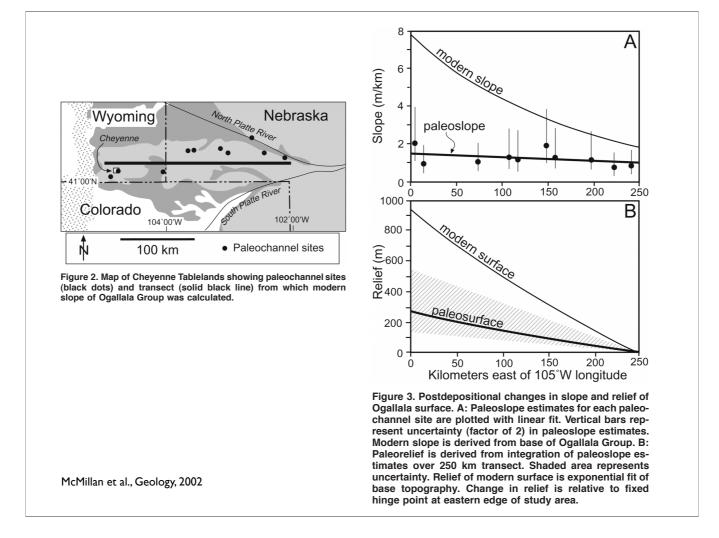
Note these are oceanic measurements...



Note these are global estimates...





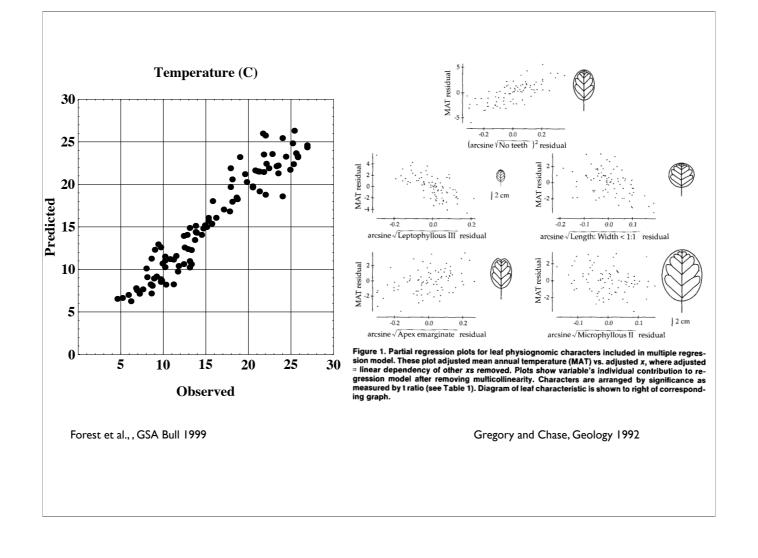


Not exactly paleoaltimeter, but use of sedimentology should tell of tilts (but watch the assumptions carefully).

What varies with elevation?

- Temperature (dry lapse rate 9.8°C/km, moist ~5°/km)
 - Paleobotany
 - Clumped isotopes
- Precipitation
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- Air pressure
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Elevation = $\frac{\text{(sea-level MAT - flora MAT)}}{\text{terrestrial lapse rate}}$ + sea level,

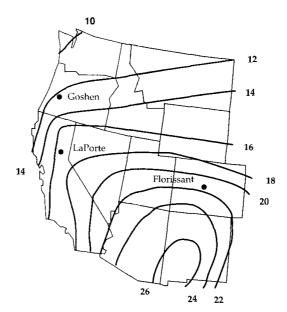
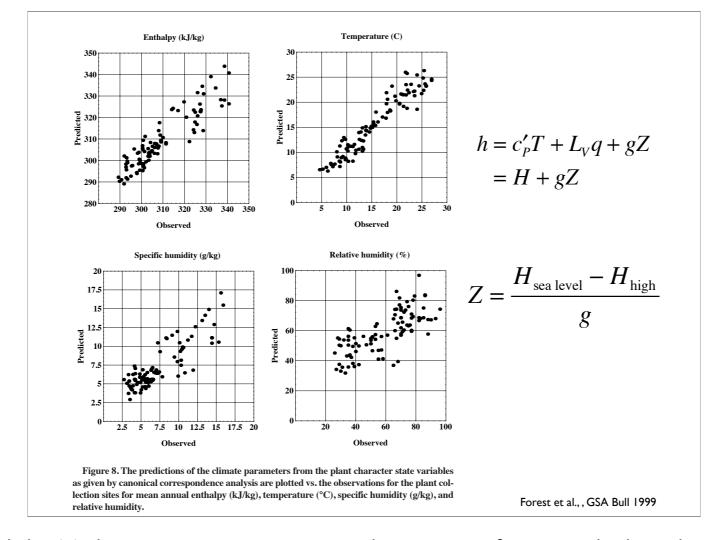
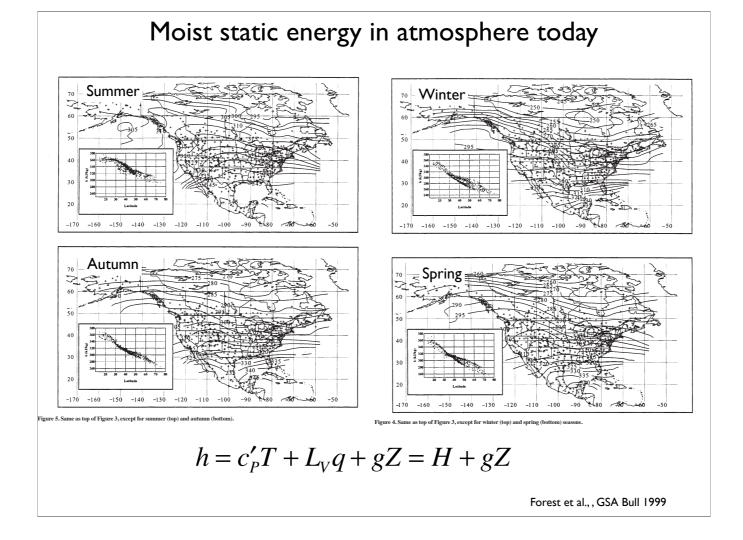


Figure 2. Mean annual temperature (MAT) at sea level in °C for modern western United States. Note that MAT increases inland due to effects of continentality and elevated base level. Redrawn after Meyer (1986).

Gregory and Chase, Geology 1992



moist static energy (h) is moist enthalpy (H) plus GPE--H is temperature times heat capacity of moist air plus latent heat of vaporization of water times specific humidity. If h is conserved, then differences in H give elevation



So is moist energy h preserved?

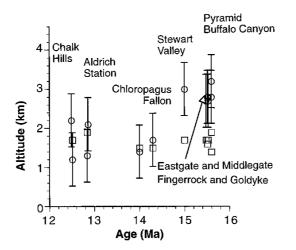


Fig. 3. Paleoaltitudinal estimates for middle Miocene leaf assemblages of western Nevada (circles) versus time. The altitude estimates for assemblages from ~15 to 16 Ma are consistently higher than present-day altitudes (denoted by squares), whereas late middle Miocene (12 to 14 Ma) assemblages have estimates that are close to present-day altitudes.



Fig. 2. Map of part of California and Nevada showing the present-day topography and the Miocene fossil sites (+) that produced the collections of leaves analyzed in this report. Numbers coordinate with those in parentheses after the assemblage names in Table 1. Not shown is Molalla, which is about 50 km southeast of Portland, Oregon, on the eastern side of the Willamette Valley.

Wolfe et al., Science 1997

What varies with elevation?

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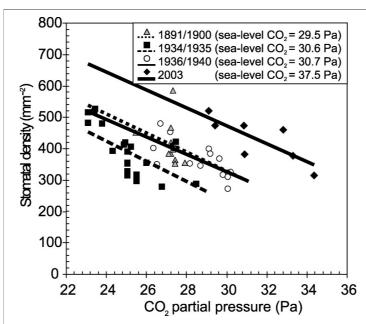


Figure 3. Historical and modern stomatal density (SD) data sets of *Quercus kelloggii* vs. pCO_2 from herbarium leaves (sun + shade) collected in (1) 1891–1900, (2) 1934–1935, (3) 1936–1940, and recent field collections, (4) 2003, demonstrating similar SD response rates—(1) -30.273, (2) -32.046, (3) -27.756, and (4) -28.528, respectively—to decreasing pCO_2 but different intercepts.

$$pp(z) = 101,325^{(-m_{\text{air}}gz/Rt)},$$

 $cd_2(z) = \frac{pp(z)}{101,325}cd_1,$

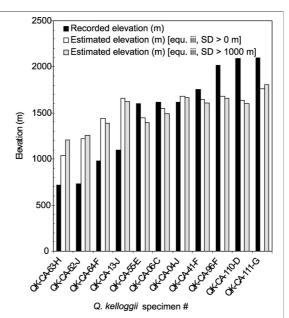


Figure 4. Actual (black bars) vs. estimated elevations of *Quercus kelloggii* collected in 2003 based on mean stomatal density (SD) of one leaf sample per tree or shrub. White and gray bars indicate elevations estimated by using equation 3, where cd_2 was estimated using linear transfer function (SD vs. pCO_2 for 1934–1935 specimens from >1000 m, Fig. 2) and nonlinear shade-leaf transfer function (SD vs. pCO_2 for 1934–1935 specimens from >0 m), respectively. Table DR2 (see footnote 1 in text) gives geographical and ecological setting of each Q. kelloggii specimen.

$$(z) = \ln\left(\frac{cd_2}{cd_1}\right) R \frac{t}{-m_{\text{air}}g},$$

McElwain, Geology 2004

pp is partial pressure of CO2 as a function of elevation, cd2

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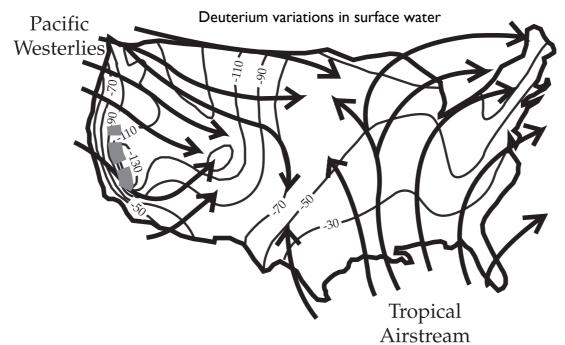


Figure 2. Map showing the principal winter season circulation patterns (arrows) and moisture sources of the United States [from *Bryson and Hare*, 1974]. Also shown are contoured δD values of surface waters and precipitation [from *Dansgaard*, 1964; *Friedman et al.*, 1964]. The approximate position of the Sierran crest is indicated by the thick dashed line. The major isotopic anomaly in eastern California and western Nevada results from the interception of westerly airstreams by the Sierra Nevada and isotopic distillation during rainout in passage across the mountains.

Poage and Chamberlain, Tectonics 2002

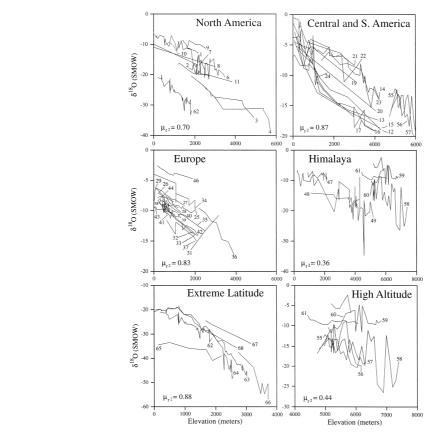
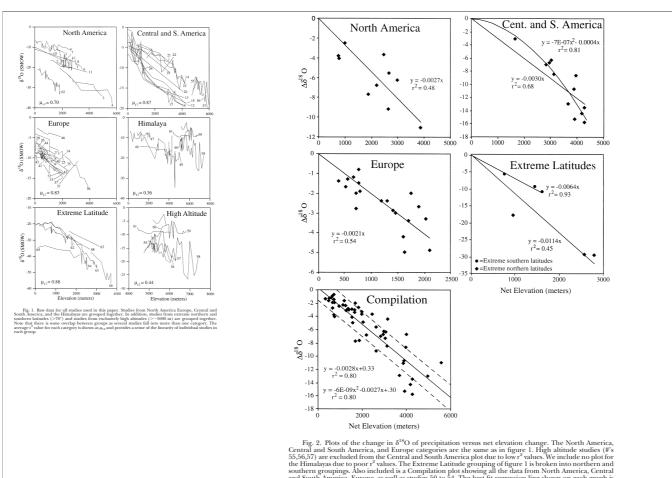


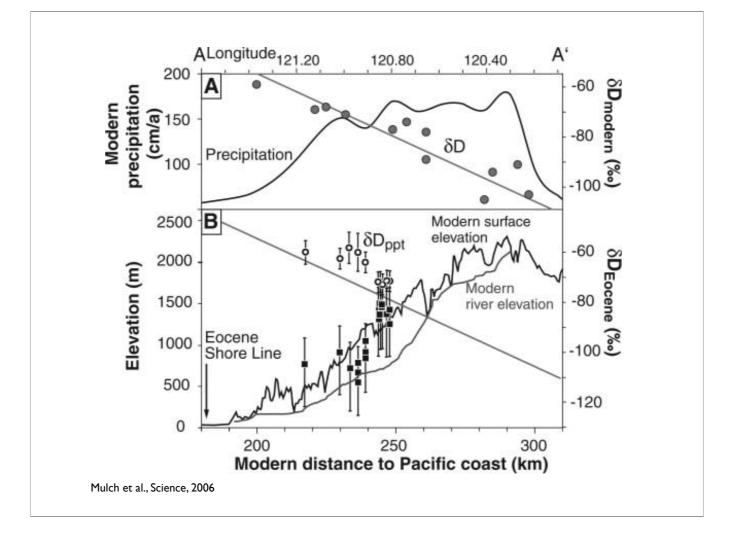
Fig. 1. Raw data for all studies used in this paper. Studies from North America Europe, Central and South America, and the Himalayas are grouped together. In addition, studies from extreme northern and southern latitudes $(>70^{10})$ and studies from exclusively high altitudes $(>70^{10})$ and studies from exclusively high altitudes $(>70^{10})$ and grouped together. Note that there is some overlap between groups as several studies fall into more than one category. The average r^2 value for each category is shown as μ_{r2} and provides a sense of the linearity of individual studies in each group.

Poage and Chamberlain, Am J Sci 2001

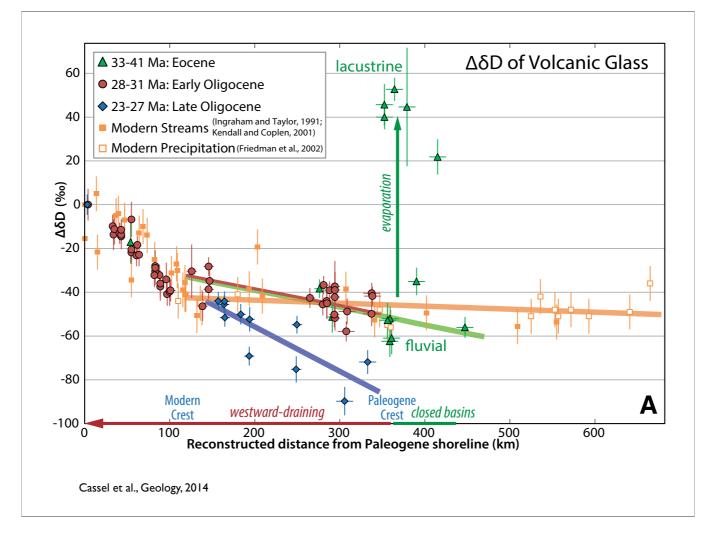


Poage and Chamberlain, Am J Sci 2001

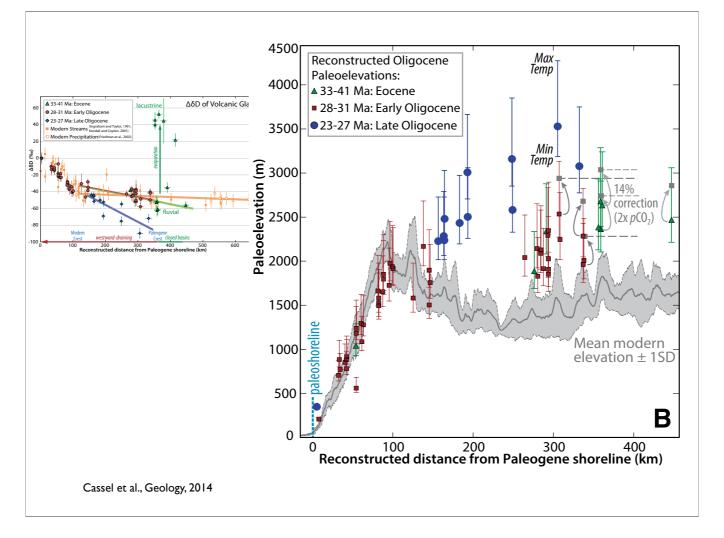
Fig. 2. Plots of the change in δ¹⁸O of precipitation versus net elevation change. The North America, Central and South America, and Europe categories are the same as in figure 1. High altitude studies (#'s 55,56,57) are excluded from the Central and South America plot due to low r' values. We include no plot for the Himalayas due to poor r' values. The Extreme Latitude grouping of figure 1 is broken into northern and southern groupings. Also included is a Compilation plot showing all the data from North America, Central and South America, Europe, as well as studies 50 to 54. The best fit regression line shown on each graph is forced through the origin except for the Compilation plot. Also shown on the Central and South America category is the best fit polynomial curve. The regression on the Compilation plot is not forced through the origin for purposes of calculating the predictive error; however, this does not change the slope of the regression line significantly. The calculated 1 standard error envelope is shown on the Compilation plot. The best fit polynomial curve for the Compilation plot is indistinguishable from the best linear fit.



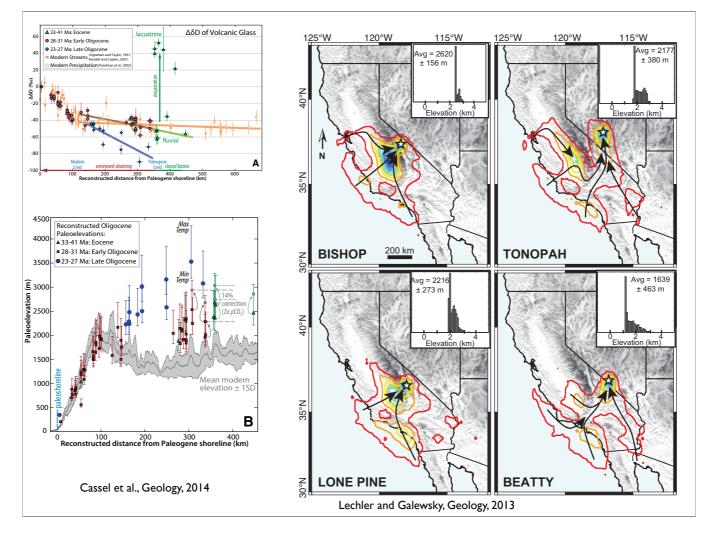
Bottom panel squares are inferred elevations from the dD values in circles.



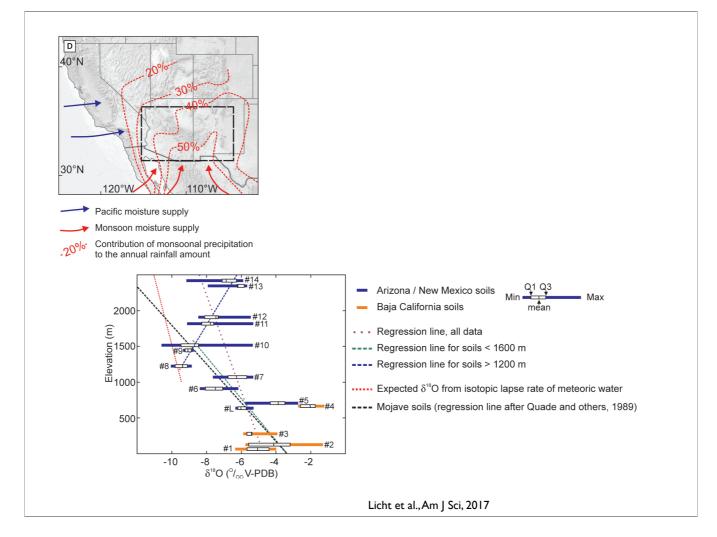
Bottom panel squares are inferred elevations from the dD values in circles.



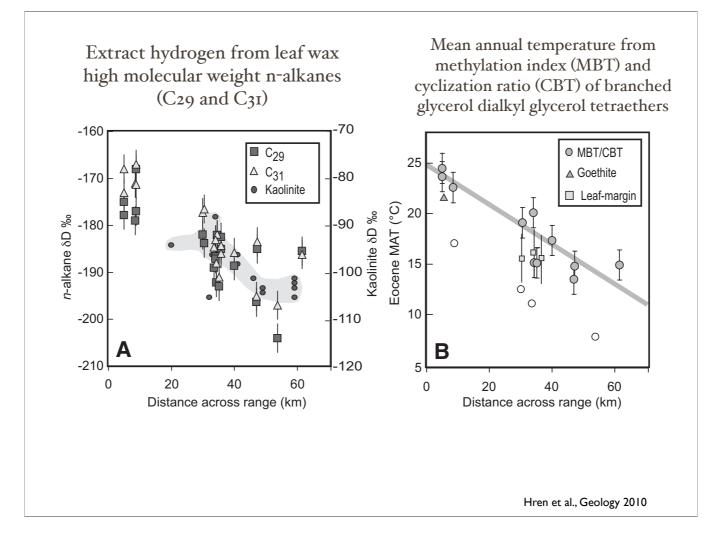
Bottom panel squares are inferred elevations from the dD values in circles.



But moisture moves around in complex ways—virtually all these proxies need a solid grounding in the climate



Monsoon in American SW can throw things into serious disarray—these are measurements from pedogenic carbonates (such as might be used for clumped isotope work).



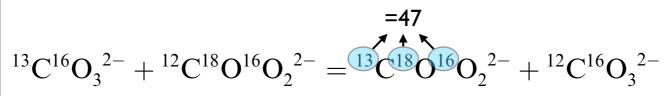
More recent work has indicated that you need to know what kind of vegetation is providing these organic molecules. There is some evidence of some fractionation in some climates as well—quite a literature blooming using this.

Paleoaltimetry

What varies with elevation?

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 - Clumped isotopes
- Precipitation
 - Paleobotany
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 - Bubbles in lava
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Favored when colder

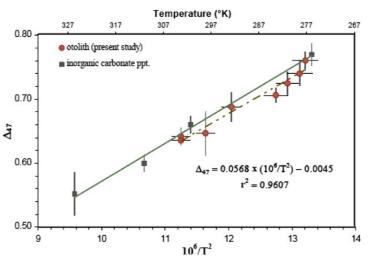
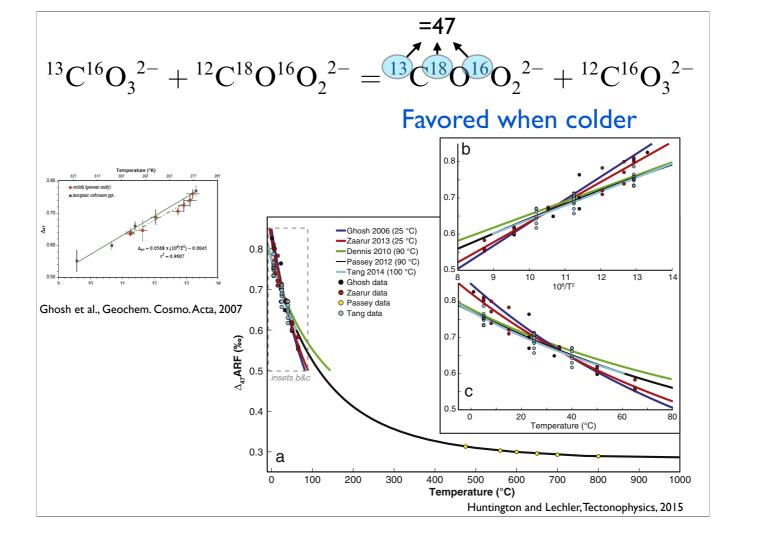
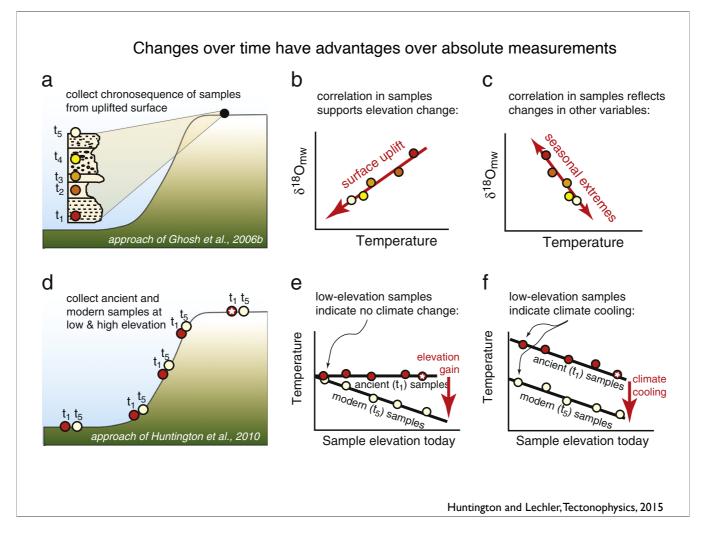


Fig. 2. Relationship between estimated environmental temperatures at which sampled fishes lived and Δ_{47} values of CO₂ produc phosphoric acid digestion of fish otoliths (plotted as filled circles). Data for inorganic calcite (plotted as filled squares) grown in the labor under controlled conditions (Ghosh et al., 2006) are shown for comparison.

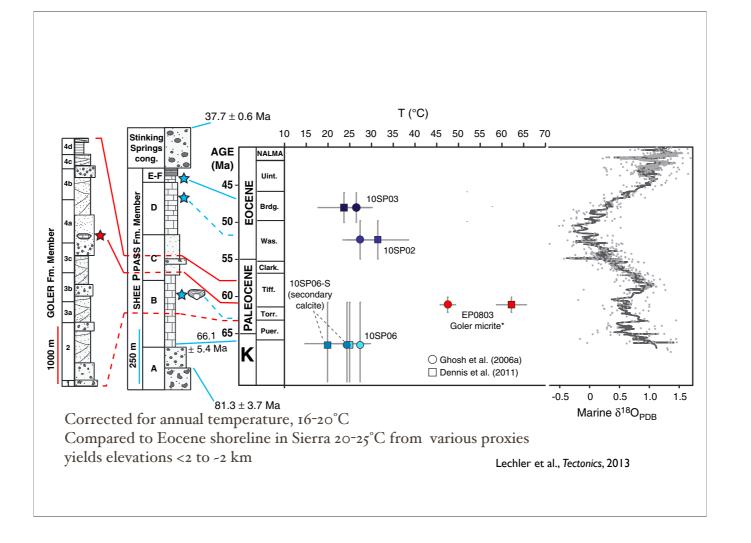
Ghosh et al., Geochem. Cosmo. Acta, 2007

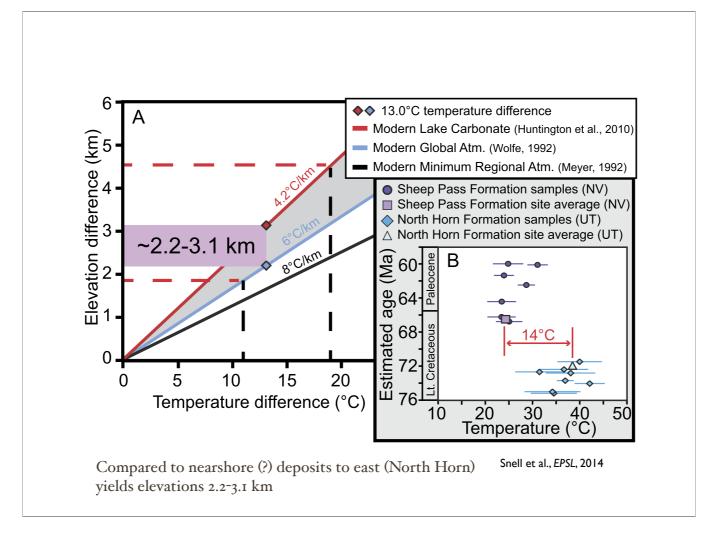


Clumped isotope approach can get temperatures and isotopes... b Street temperature Lower 8¹⁸O_{mw} Higher 8¹⁸O_{mw} Higher 8¹⁸O_{mw} Higher 8¹⁸O_{mw} Higher 8²⁸H_{mw} Higher 8²⁸H_{mw}

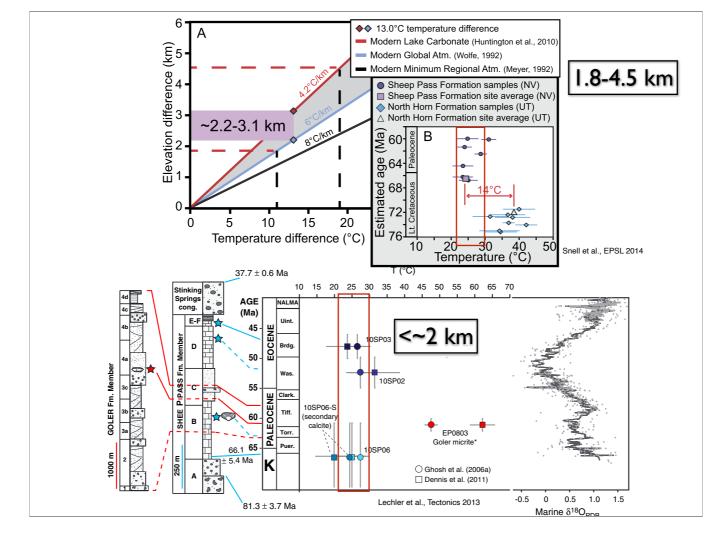


The problems with clumped isotopes revolve around precipitation of the carbonate (seasonal? which season?) and diagenesis





Now you could say these agree at 2 km, but each paper argues away from that number. But note the raw temperatures are



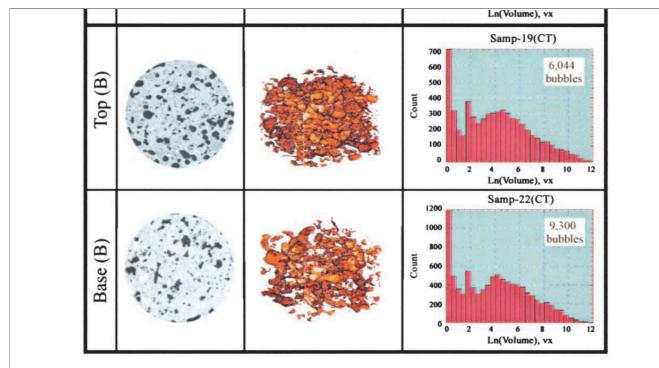
OK, so why the two different results. Lechler et al. then correct these values, which are assumed to be summer season, to a mean annual temperature of 15-20C. They then compare with coastal CA estimates of 20-25C. Snell et al. compare North Horn and Sheep Pass directly

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	Thickness	N bubbles	Modal size mm ³	Pressure (atm)	Inferred Elevation (m)	Actual Elevation (m)
Тор	1.65	6044	1.823	0.704	3254	3932
Base		9300	1.125		:	Sahagian et al., J.

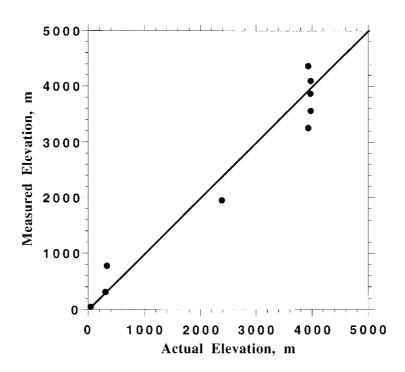


Figure 9. Results of analysis of using vesicular basalt as a measure of paleoelevation. Perfect results would lie along the diagonal line.

Sahagian et al., J. Geol., 2002

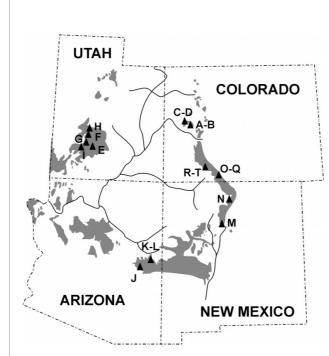


Figure 1. Sampling localities throughout perimeter of Colorado Plateau. Lava fields are shaded. In general, southern localities are limited to relatively young flows, while older flows can be found in northern sections. Oldest flows are andesitic. There is no relationship between paleoelevation and age (see Fig. 2).

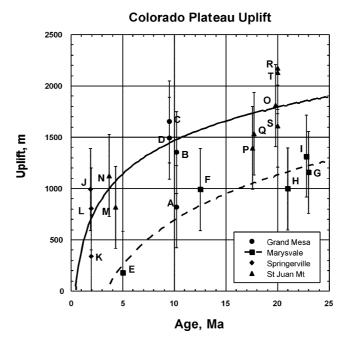
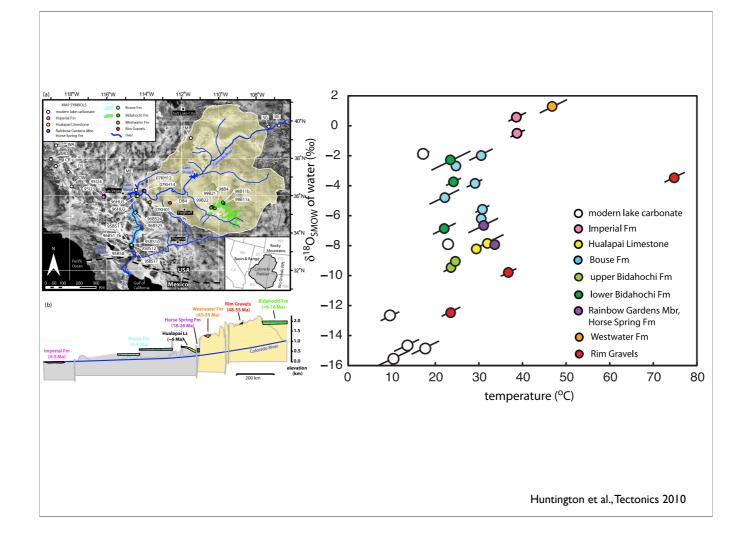
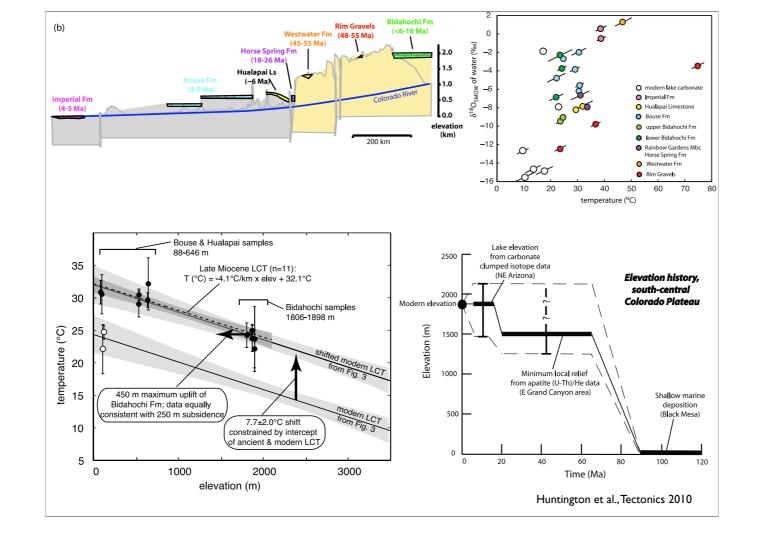


Figure 3. Uplift history of Colorado Plateau based on vesicular basalt paleoaltimeter. Logarithmic curve provides highest r-value (solid curve) relative to other curve fits. However, a logarithm passes through (0,0) even though this point is actual value of "no uplift since present." Results indicate that slow uplift (40 m/m.y.) commenced at least 25 m.y. ago, but accelerated (to 220 m/m.y.) in past 5 m.y. Dashed line is for Marysvale samples collected from downfaulted blocks of transition zone to Basin and Range.

Sahagian et al., Geology 2002





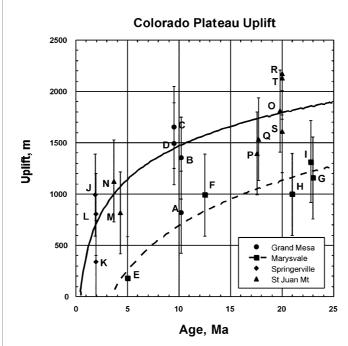
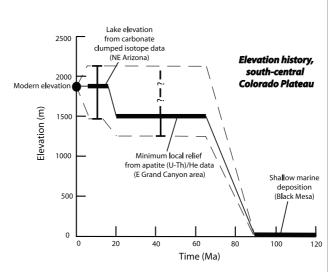


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Sahagian et al., Geology 2002



Huntington et al., Tectonics 2010